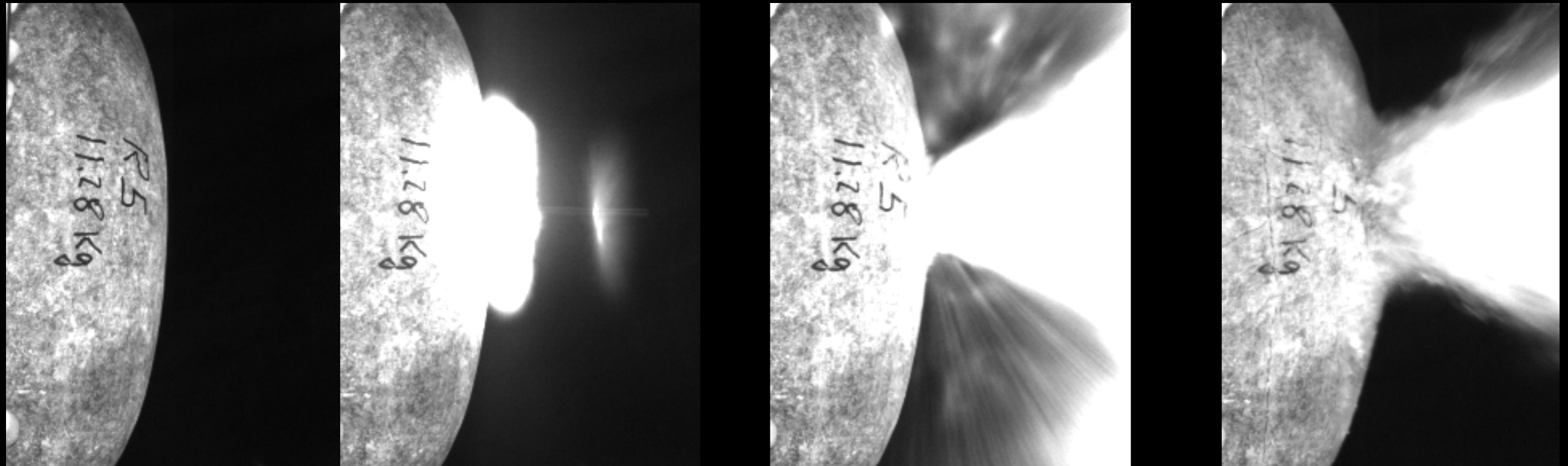


# Momentum transfer via direct impact: Experimental measurements



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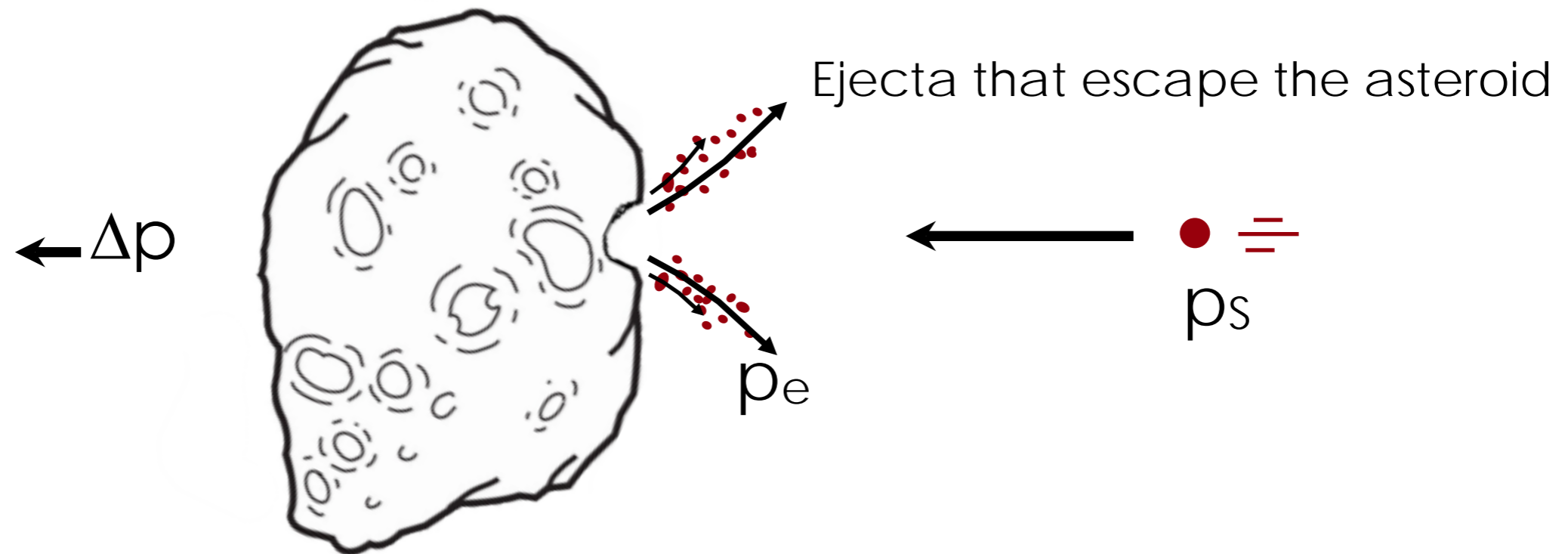
# The central question

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- i How effective is the direct impact method?
  - What is the largest asteroid we could expect to deflect?



# Momentum multiplication factor



Change in asteroid momentum:  $\Delta p = p_s + p_e$

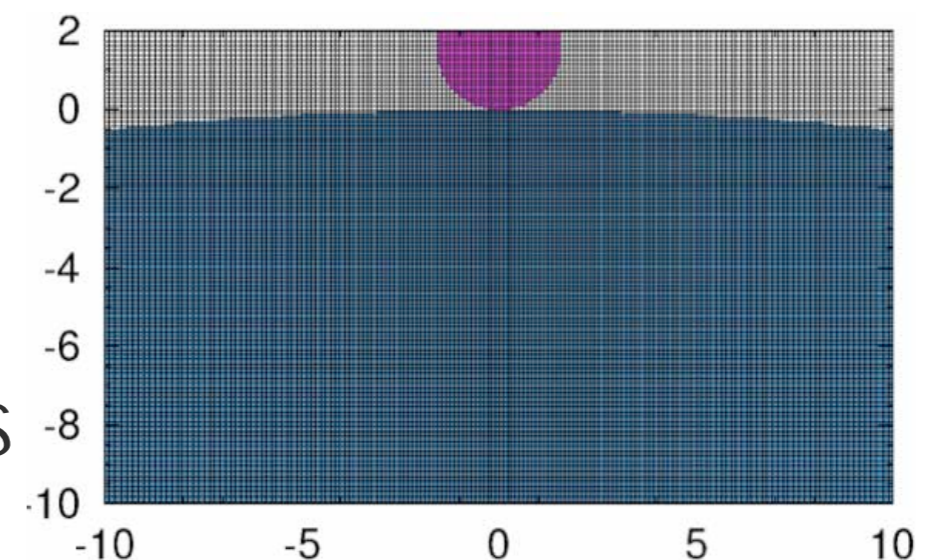
Define  $\beta = \frac{\Delta p}{p_s}$

$$\beta = 1 + \frac{p_e}{p_s}$$

$\beta > 1$ , unless there is no escaping ejecta

# Our approach

- i Conduct impact experiments to directly measure  $\beta$ .
  - But can't do experiments under the conditions of an actual mission.
  - So, the experiments are “scaled” to the mission conditions.
- i Numerical simulations
  - Validate against experiments
  - Extend the reach of experiments



# Scaling of ejecta momentum

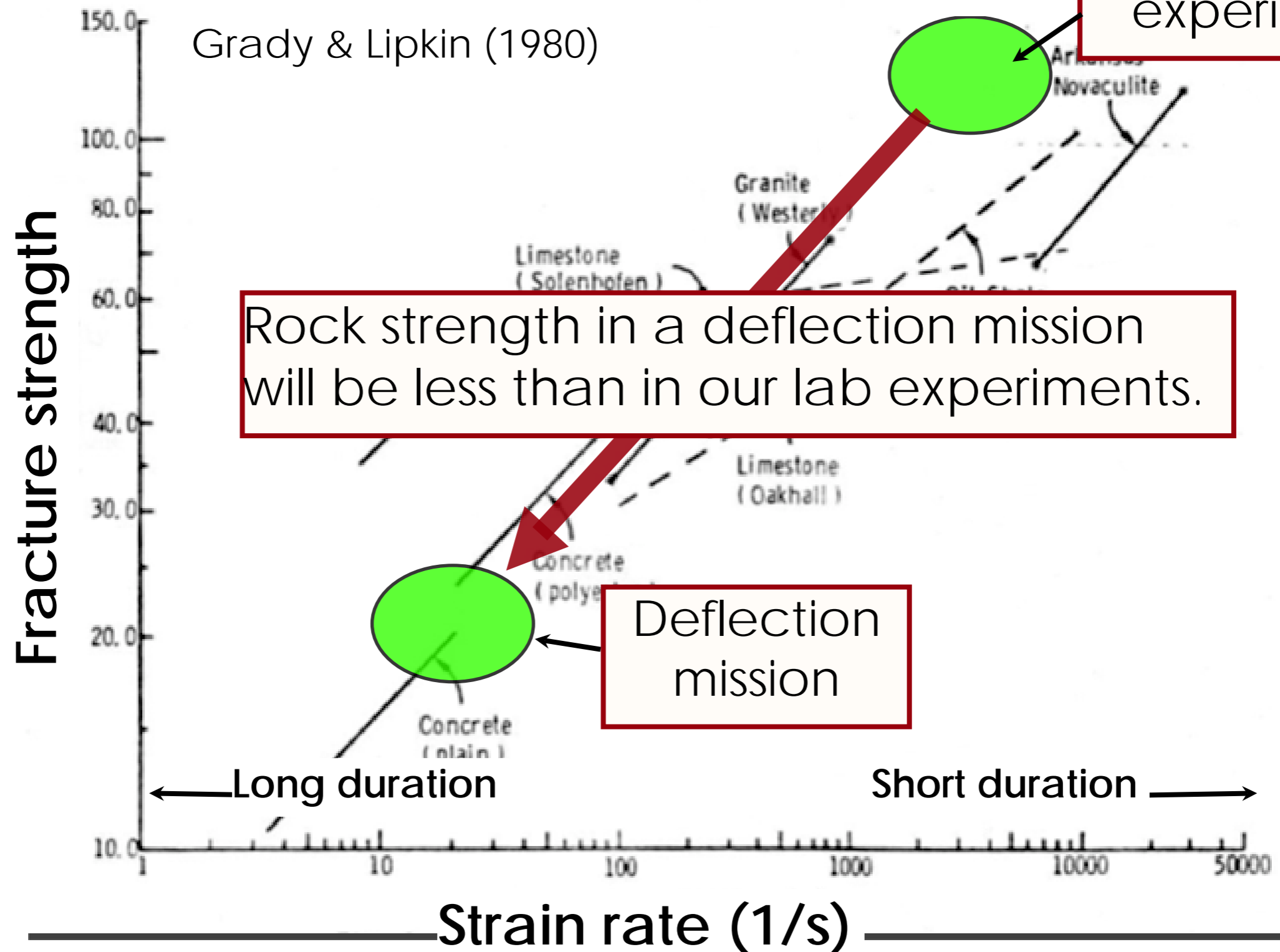
$$\beta = 1 + \frac{p_e}{p_s}$$

- i Understanding  $\beta$  is just a cratering and ejecta problem.
- i We know how ejecta scales from point-source theory of cratering.
  - (Housen et al 1983, Holsapple & Schmidt 1987, Housen & Holsapple 2011)

# Two regimes of crater formation

- i Gravity regime  $\rho gh \gg \text{cohesion}$ 
  - Large-scale impacts, or craters formed in very low strength materials (e.g. sand)
- i Strength regime  $\rho gh \ll \text{cohesion}$ 
  - “Small” impacts in rocky targets
  - And perhaps “cohesionless” soils at micro-G
- i Applying lab results to a mission requires knowing which regime you’re in.

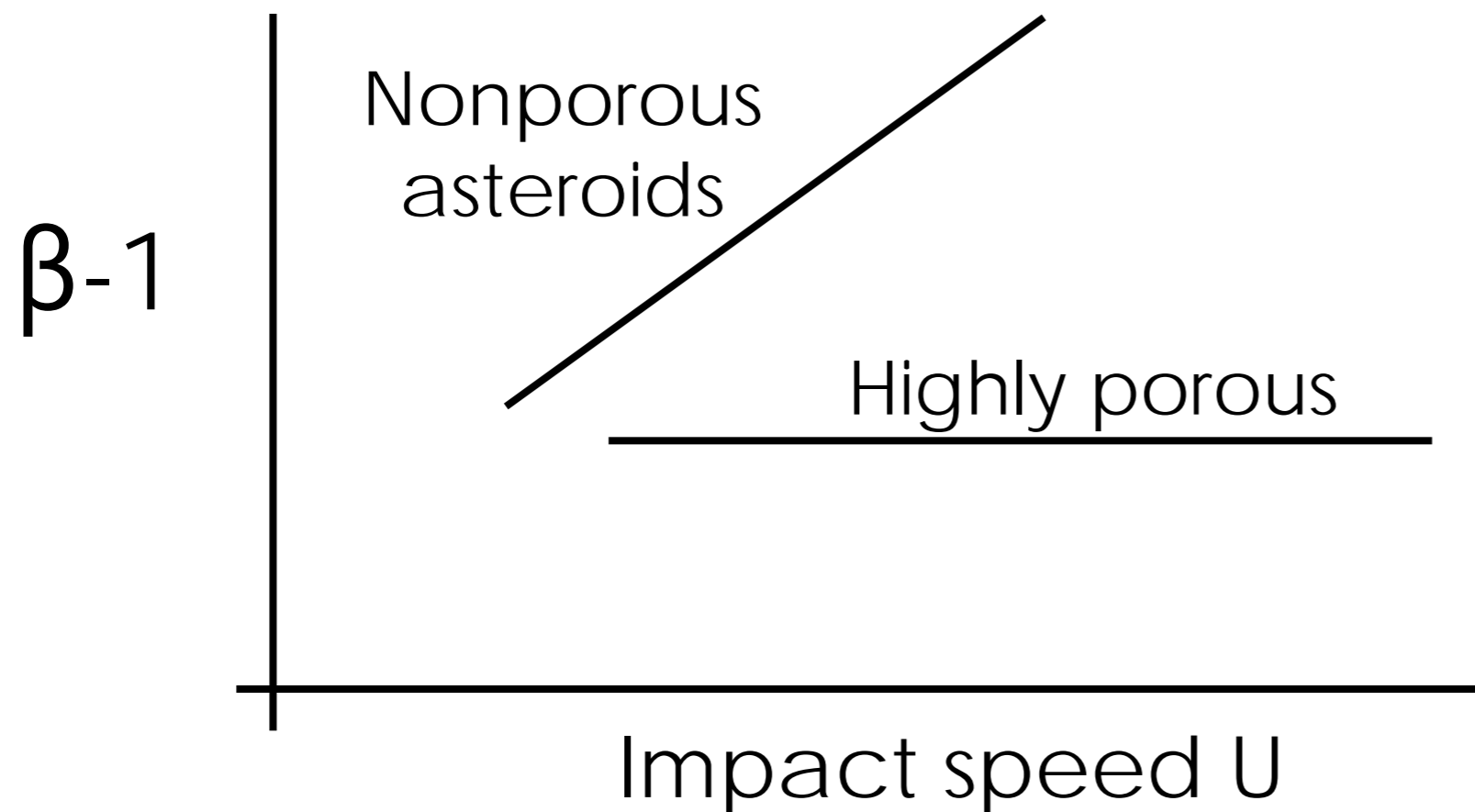
# Rock strength is rate dependent



# Scaling of $\beta$

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	Nonporous	Moderate porosity	Highly porous
Strength	$\beta - 1 \propto a^{0.15} U^{0.73}$	$\beta - 1 \propto U^{0.2}$	$\beta - 1 \approx \text{constant}$
Gravity	$\beta - 1 \propto (ga)^{-0.25} U^{0.5}$	$\beta - 1 \propto (ga)^{-0.08} U^{0.16}$	$\beta - 1 \approx \text{constant}$



Low-porosity materials have the strongest velocity dependence

# Experiments to measure $\beta$

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- i Ballistic pendulum
  - Works, but granular materials need to be shot vertically
- i Force transducers
  - Uncertainty when integrating force-time history to get impulse
- i Suspend target from springs



Impact chamber in Boeing Shock Physics Lab

# Experiments to measure $\beta$

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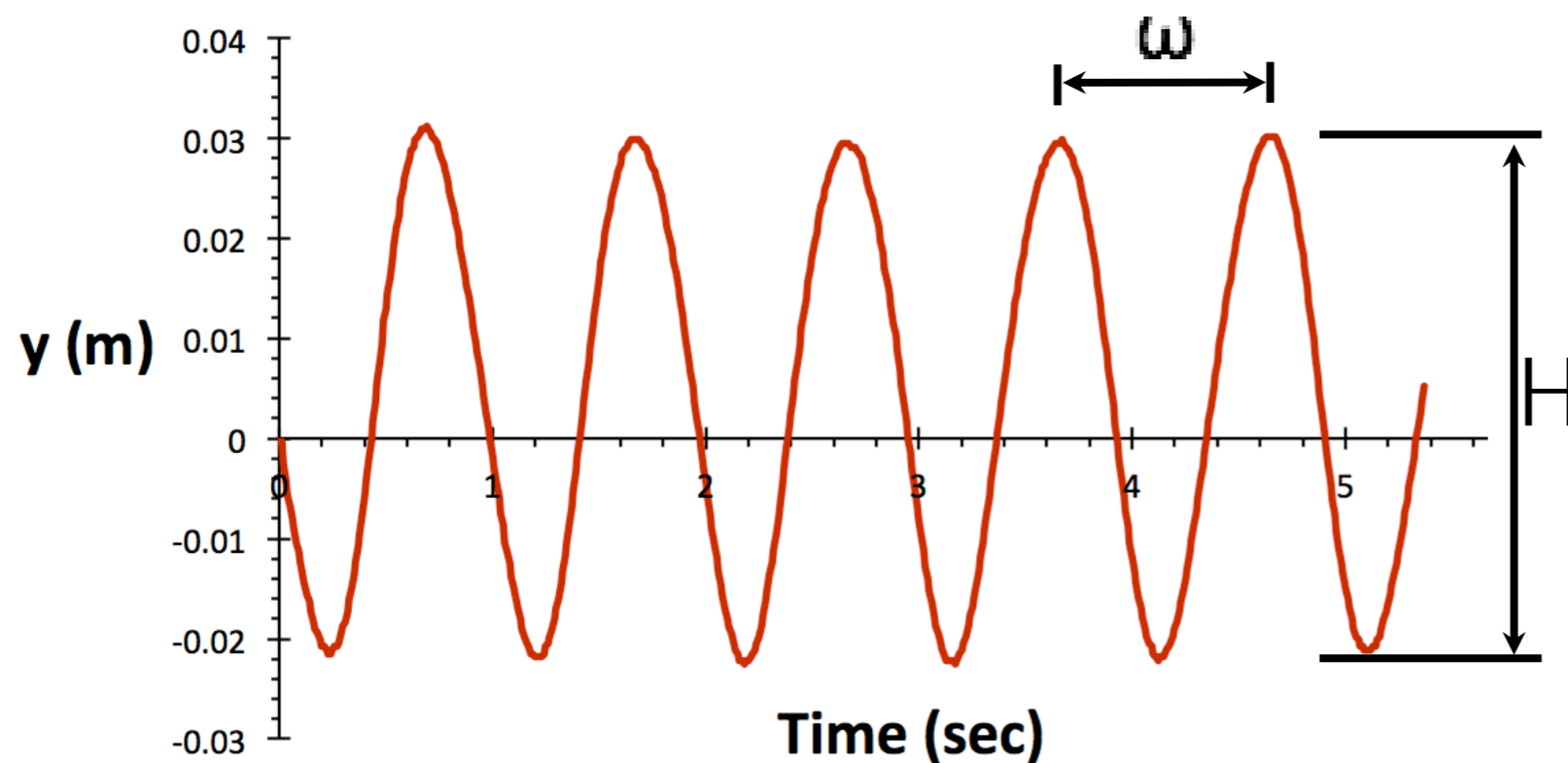
NASA Ames Vertical Gun Range



# How we measure $\beta$

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Sand target  
5.7 km/s impact  
6.4 mm Al sphere



Impulse delivered to target:

$$I = \frac{H\omega M}{2}$$

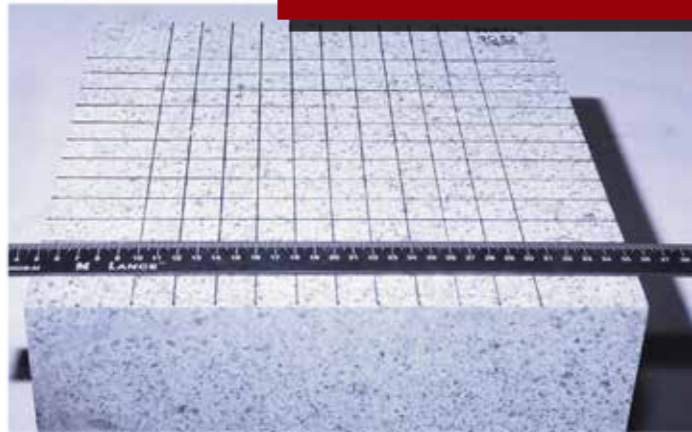
$$\beta = \frac{I}{mU}$$

$m, U$  = projectile mass, speed

# Target materials used to-date

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Low porosity  
High cohesion



River rock

Granite & basalt blocks

High porosity  
Medium cohesion



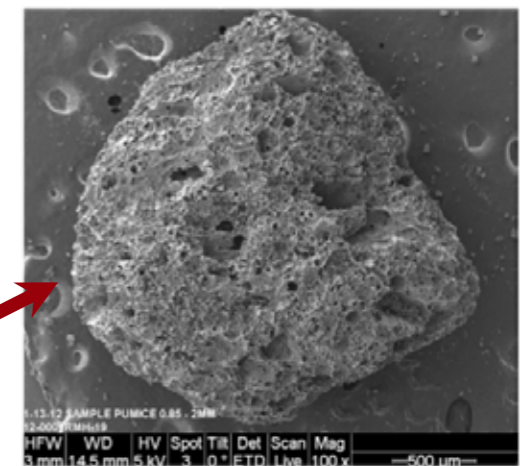
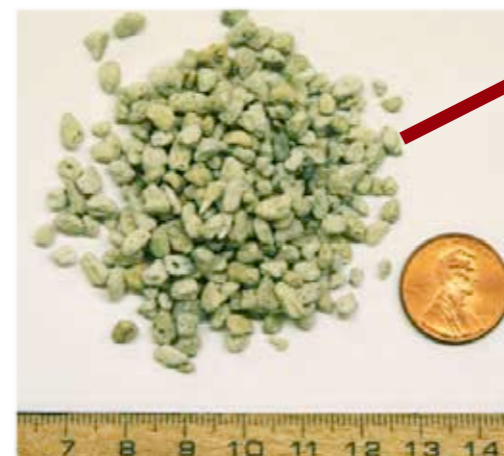
Cohesive  
pumice

Medium porosity  
No cohesion



Dry sand  
35% porosity

High porosity  
No cohesion



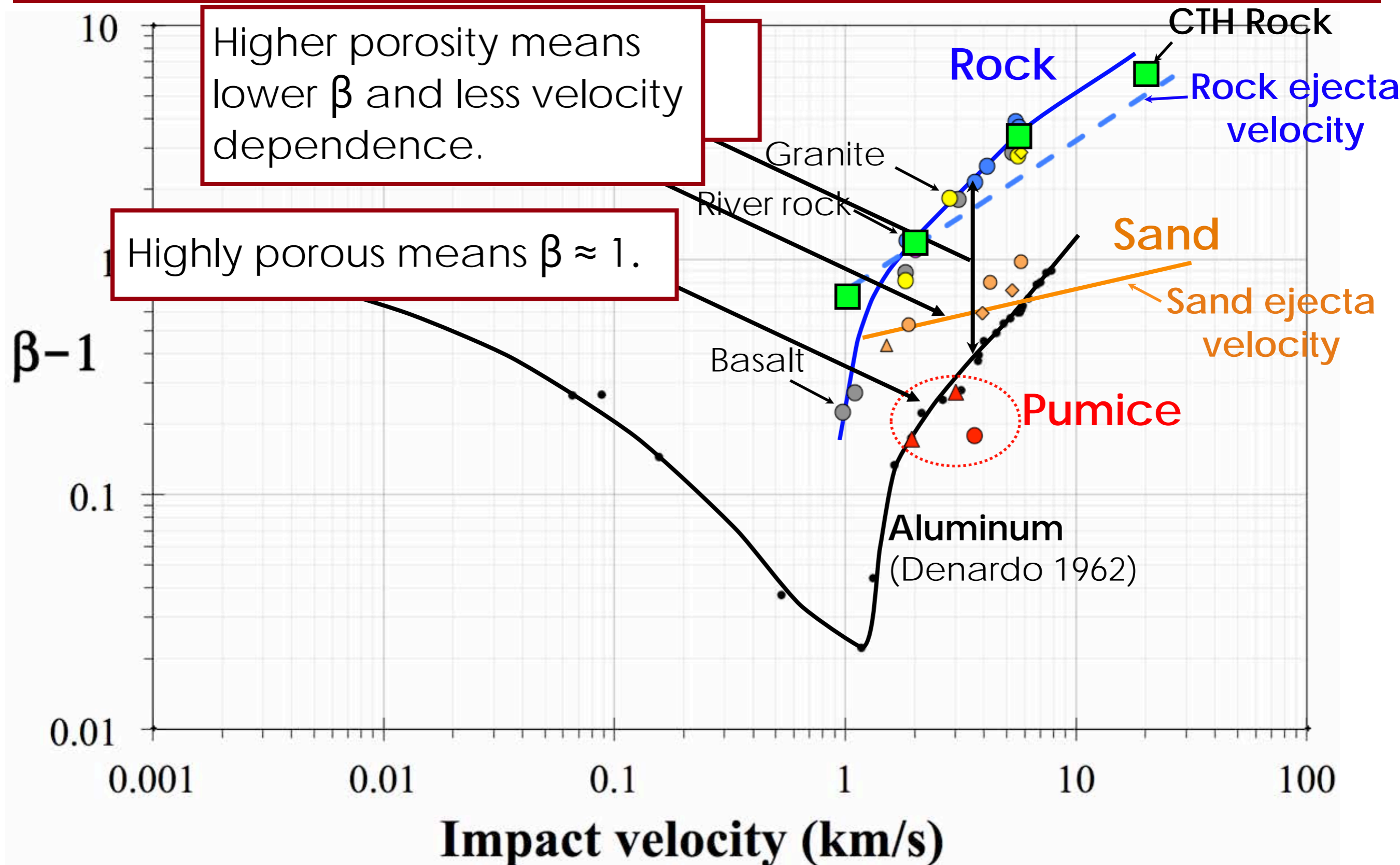
Granular pumice  
84% porosity

# Results

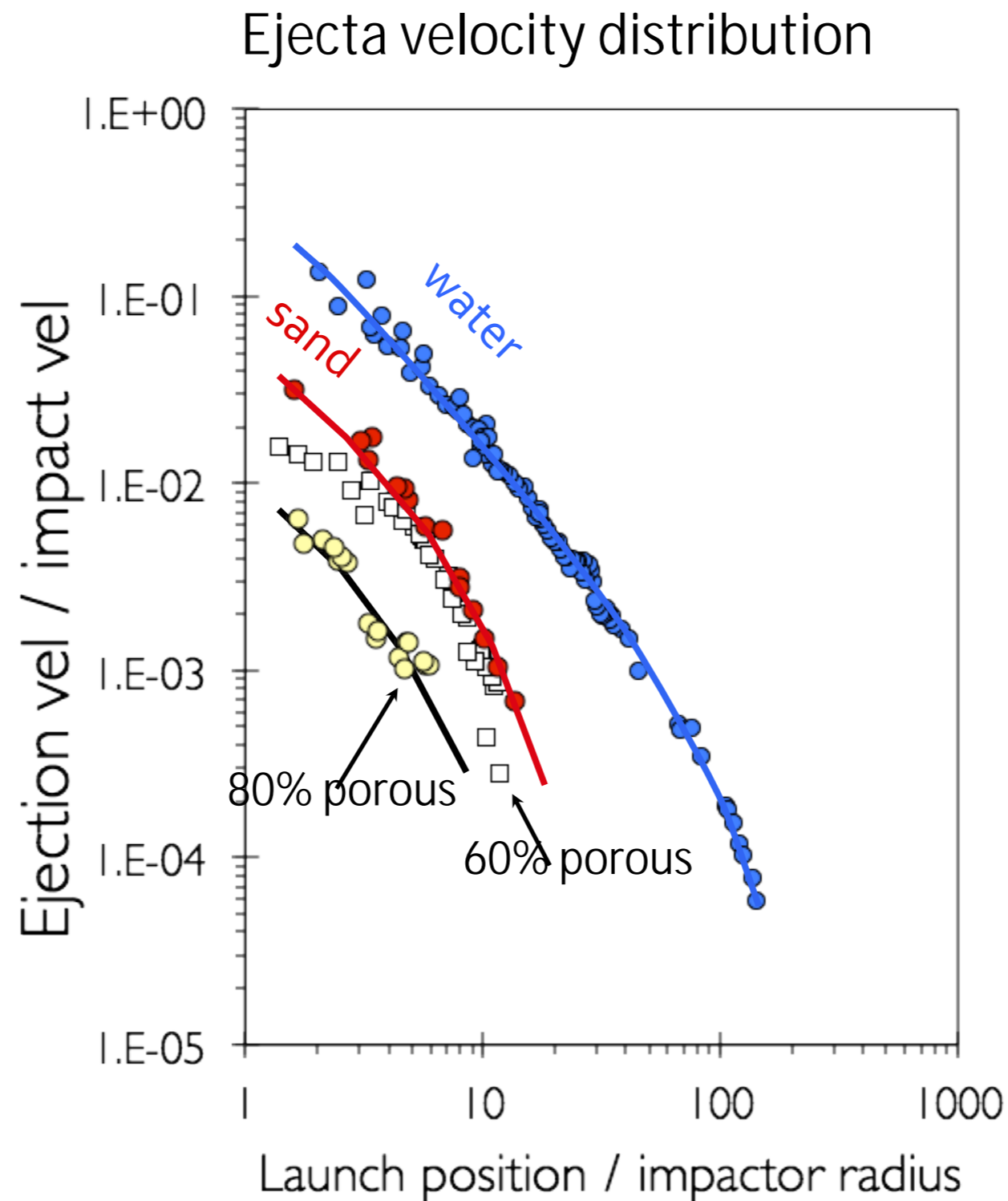
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Higher porosity means lower  $\beta$  and less velocity dependence.

Highly porous means  $\beta \approx 1$ .



# High porosity reduces ejecta velocity



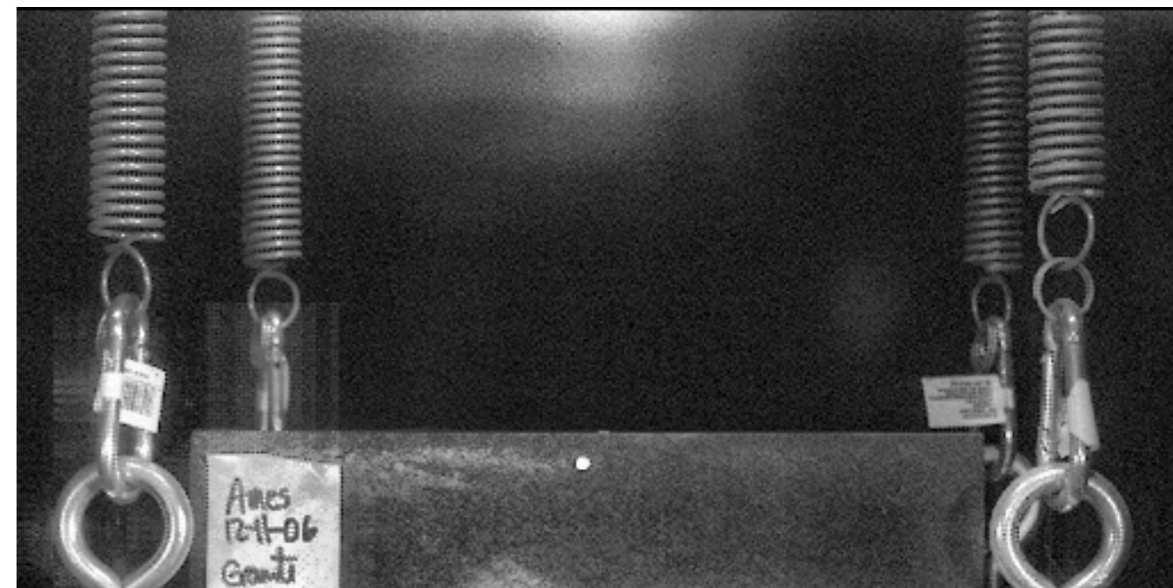
- i Is the event strength or gravity dominated?
  - The crater scaling is different in these two regimes.
- i How much of the ejecta escapes the asteroid, and how much momentum does it contain?

# Rocky targets

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- i Strength dominated at all size scales of interest.
- i All ejecta would escape an asteroid.
  - Slowest ejecta has m/s speeds.
  - Escape speed for 500 m asteroid = 0.2 m/s

Granite target  
5 km/s impact  
6.4 mm Al sphere



# Scaling to higher impact speed

5 km/s lab experiment



Measured  
 $\beta \approx 4 - 5$

20 km/s mission



4x higher  
velocity

$$\beta - 1 \propto U^{0.73}$$

$$\beta = 9-12$$

CTH @ 20 km/s:  $\beta = 7.5$

Conservative:

$$\beta \approx 8$$

# Strain rate effect

$$\beta - 1 \propto a^{0.15}$$

↙ Impactor size

- i Strength of rock depends on strain rate.
  - Rock is weaker for large-scale impacts.
  - Lower strength means bigger craters, more ejecta.
- i For a 1-m impactor, we get a factor of 2x increase in  $\beta - 1$ .

20 km/s mission to a rocky asteroid:

$$\beta \approx 15$$

# Implications for deflection

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$\Delta V = 0.01$  m/s for 1-decade warning  
(Ahrens & Harris, 1992)

$$\beta = \frac{M\Delta V}{mU} = 15$$

5-ton spacecraft, 20 km/s

A 5-ton impactor at 20 km/s imparts  $\Delta V \geq 0.01$  m/s for rocky asteroids up to ~500m diameter.

1-ton spacecraft could deflect a ~300m diameter body.

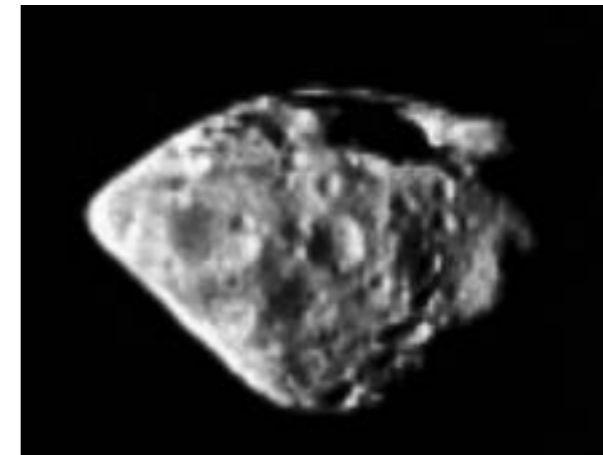
Additional conservatism - assume no rate dependence:  
5-ton spacecraft could deflect a ~300m diameter body.

- i Lab experiments are gravity dominated.
- i Assume gravity dominated for a small asteroid as a bounding case.
- i Must account for ejecta that are retained.
  - Holsapple and Housen (2012) Icarus 221.
- i Bottom line:
  - $\beta \approx 1$  to 2 for a 20 km/s deflection mission.
  - Deflection of 200 to 300m bodies is possible.

# Regolith-covered rock

- i All asteroids imaged to date show signs of a regolith-covered surface.
  - So bare rock may be an unlikely case.
- i What is the effect of a regolith layer?
- i Impacted basalt target covered by a layer of sand
  - Only have 1 experiment.

Asteroid Steins ~5 km dia



- i We find significant momentum multiplication for rocky targets, or regolith-covered rock.
- i Increased target porosity causes
  - dramatic decrease in  $\beta$  as target porosity increases
  - weaker dependence on impact speed
- i For a decade of warning deflection is possible for
  - Rocky bodies up to ~500 m diameter
  - “Balls of sand” up to 200-300 m diameter

# Open questions

- i Effect of a regolith layer
  - How does the layer thickness affect  $\beta$ ?
  - What is the impact velocity dependence?
- i Strain rate effects should be evaluated directly in momentum transfer experiments with rock targets.
- i How does the “fabric” of the target affect  $\beta$ ?
- i What is the cohesive strength of granular materials at  $\mu$ -G gravity?
  - How does that affect  $\beta$ ?